



Tracing Time Scales of Fluid Residence and Migration in the Crust

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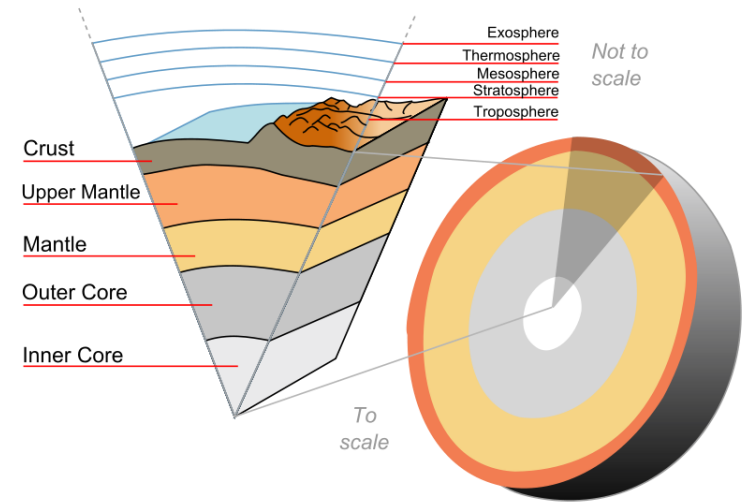
THE UNIVERSITY OF
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UIC u^b

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BERN



The Crust

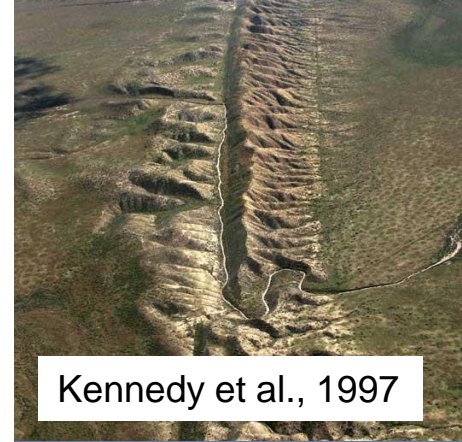


- The Crust?
 - Every naturally occurring solid you can see on Earth
 - 5-50km thick from the surface
 - Geological extract containing high concentration of U, Th, K (Production of ^4He , ^{21}Ne , ^{39}Ar , ^{40}Ar , ^{85}Kr , $^{131-136}\text{Xe}$)
- Fluids in the crust?
 - Groundwater (H_2O)
 - Natural gases/Petroleum (Hydrocarbons)
 - Volatiles associated with magmatism (CO_2 -dominated)

Fluids in the crust

Evidence of interaction with the crust

- Gradual addition of radiogenic/nucleogenic isotopes as flow distance increases



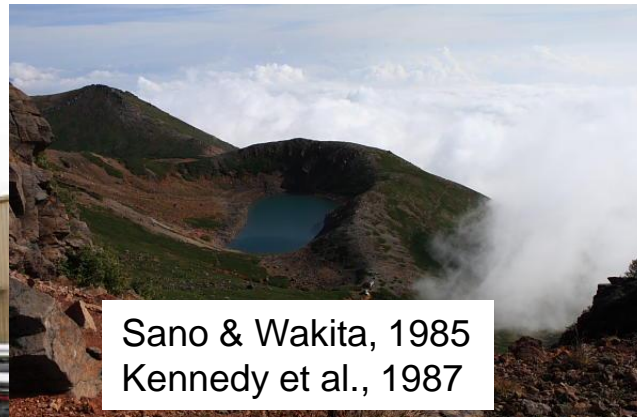
Kennedy et al., 1997



Torgersen et al., 1989
Lehmann et al., 2003



Ballentine et al., 2005
Gilfillan et al., 2008; 2009



Sano & Wakita, 1985
Kennedy et al., 1987



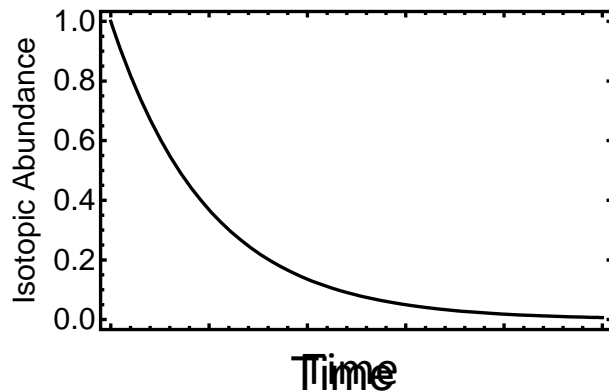
Pinti and Marty, 1995

Geochemical system in the crust

Initial

Beauty of
Noble gas radionuclides:
Simplicity

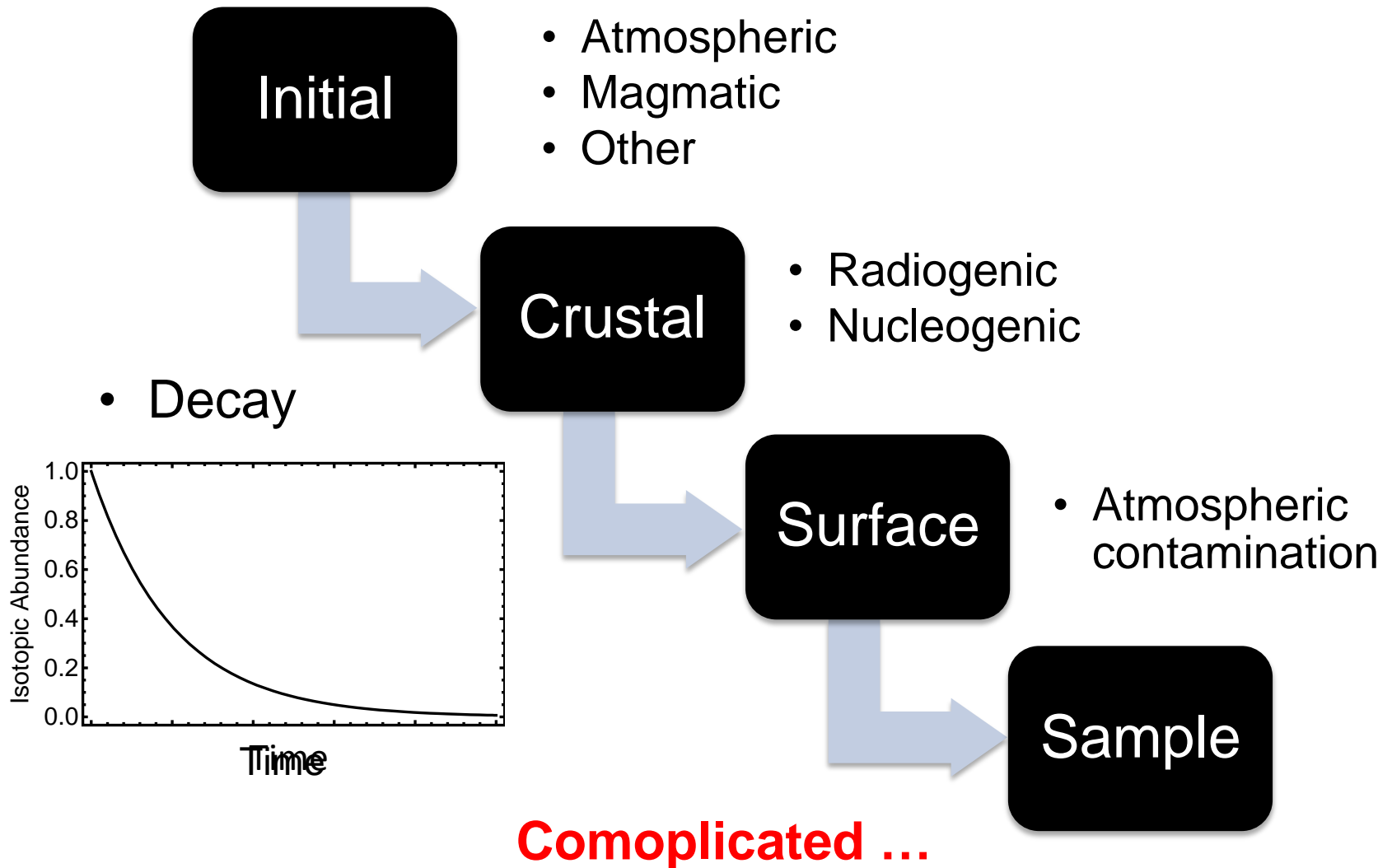
- Decay



Comoplicated ...

Sample

Geochemical system in the crust



Outline

1. Introduction: Fluids in the Crust

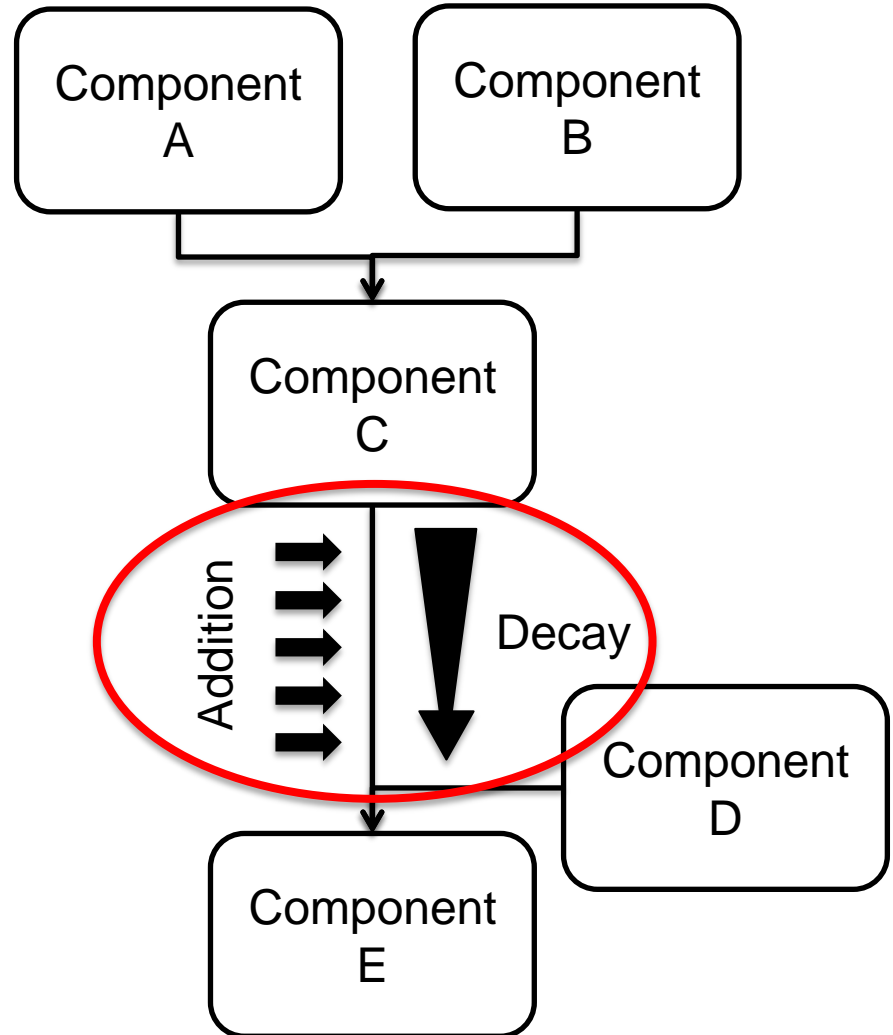
2. Theoretical analysis of complex geochemical system

1. Case study at Yellowstone National Park

Complex Geochemical System?

Processes

- Mixing
 - Episodic:
 - Magmatic volatiles
 - Atmosphere at different stages



Continuous addition of radioactive isotopes

Model:

Loss by decay

$$\frac{d^i C_{Tot}}{dt} = -\lambda_i \times {}^i C_{Tot} + \frac{d^i C_R}{dt}$$

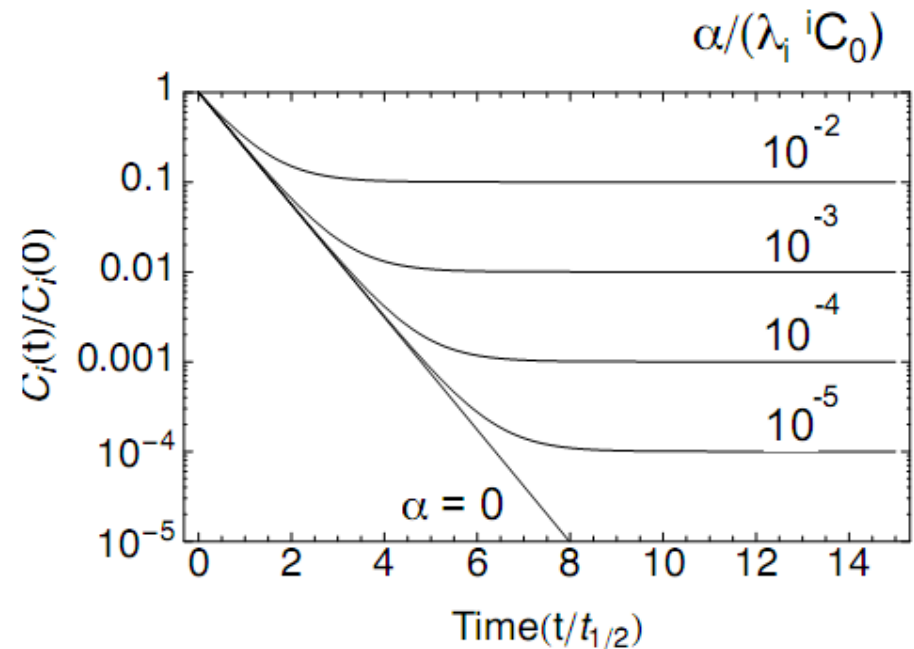
Addition from the Crust

For $d^i C_R/dt = \alpha$ (constant):

$${}^i C(t) = {}^i C_0 \times e^{-\lambda_i t} + \frac{\alpha}{\lambda_i} (1 - e^{-\lambda_i t})$$

Supply-decay equilibrium

Modeled Isotopic abundance



Notes:

- Input rate determined once in equilibrium (α/λ)
- No direct implication on "Age"

Continuous addition of stable isotopes

Model:

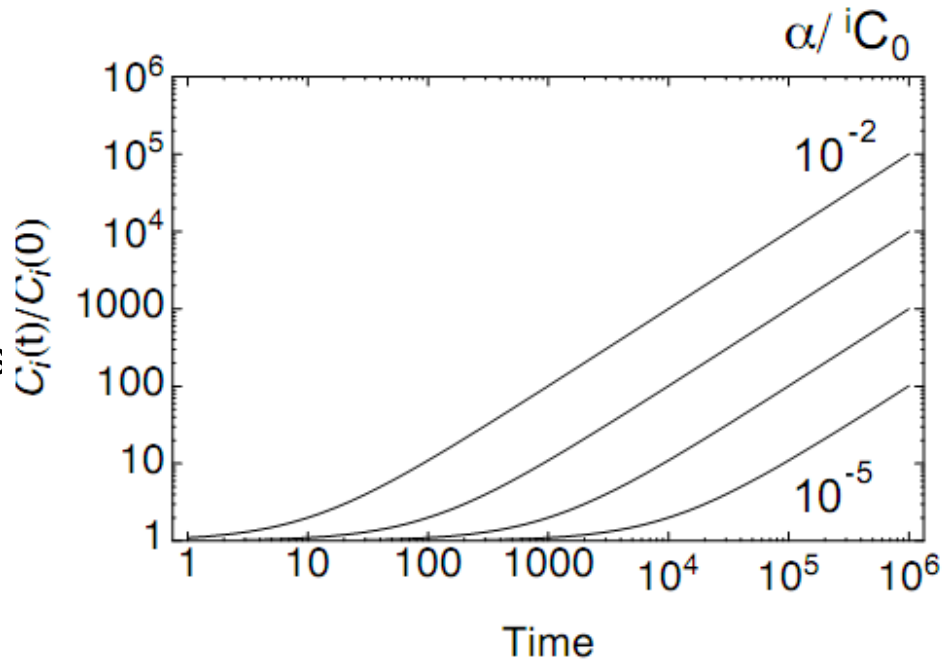
$$\frac{d^i C_{Tot}}{dt} = -\lambda^i C_{Tot} + \frac{d^i C_{\mathbb{R}}}{dt}$$

Addition from the Crus

For $d^i C_{\mathbb{R}}/dt = \alpha$ (constant):

$$^i C(t) = ^i C_0 + \alpha \times t$$

Modeled Isotopic abundance



Notes:

- Simple increase with time
- Input rate is unknown (assumed sometimes for ^4He)

Evolution of isotope ratios

Model

$${}^{ij}R_F(t) = \frac{{}^iC(t)}{{}^jC(t)} = \frac{{}^iC_0 \cdot e^{-\lambda_i t} + \frac{a_i}{\lambda_i} (1 - e^{-\lambda_i t})}{{}^jC_0 + a_j \cdot t}$$

$$\rightarrow \frac{a_i / \lambda_i}{{}^jC_0 + a_j \cdot t}$$

Steady-state

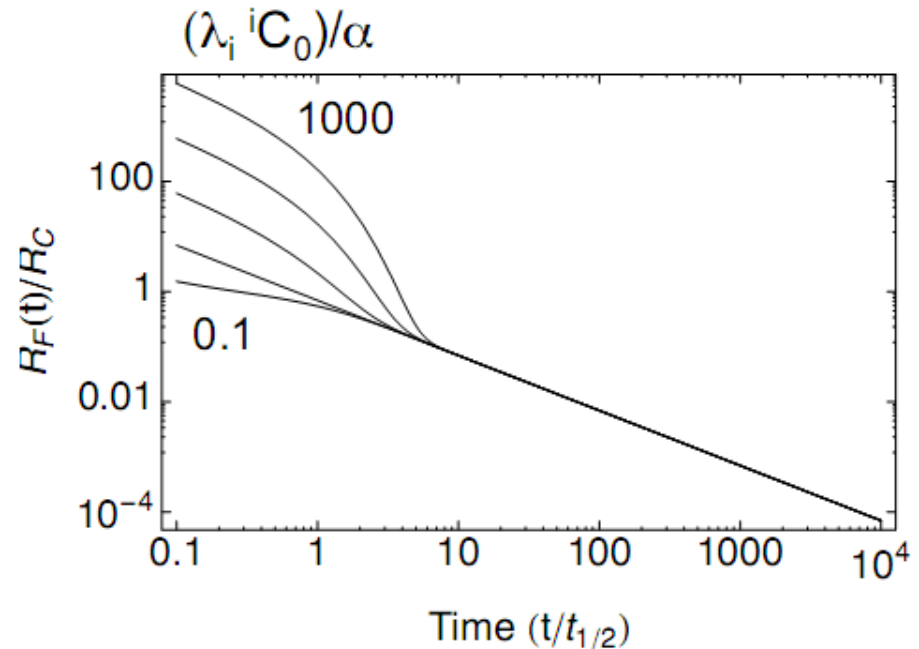
$$\rightarrow \frac{a_i / a_j}{\lambda_i \cdot t}$$

Initial correction

$$\rightarrow \frac{{}^{ij}R_C}{\lambda_i \cdot t}$$

Common α

Modeled isotope ratios



Notes:

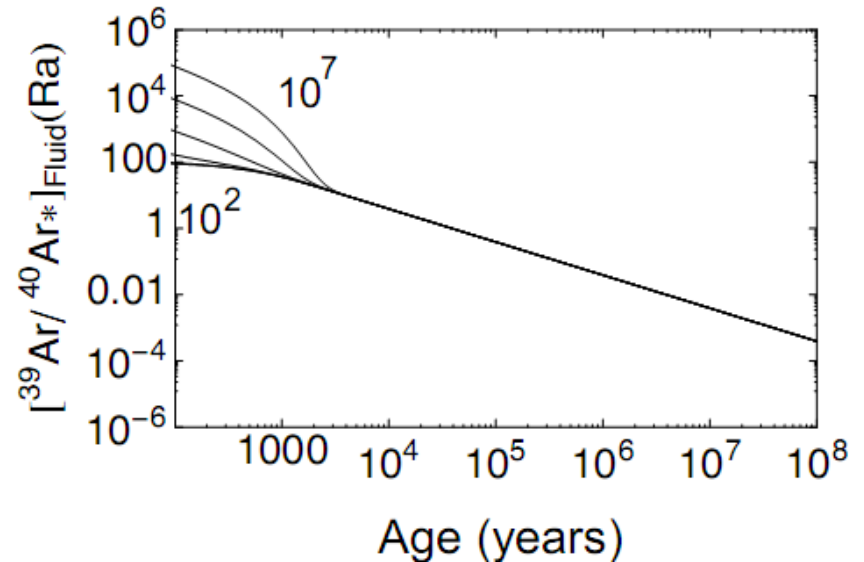
- Input rate from radionuclide
- Accumulation from stable isotope
- **New Chronometer!**

Applicable system

Isotopes

- Radioactive isotopes produced in-situ
 - ^{39}Ar by $^{39}\text{K}(n,p)^{39}\text{Ar}$
 - ^{85}Kr by ^{238}U fission
 - No significant ^{81}Kr
- Stable isotope
 - ^4He and ^{40}Ar
 - Concentrations of other non-radiogenic stable isotopes in rocks are low
- Similar α value: $^{39}\text{Ar}/^{40}\text{Ar}$

An example:



- Ar input from a reservoir rock with $^{39}\text{Ar}/^{40}\text{Ar}^*$ ratio of 100 Ra.
- The ratio evolves independent of the input rate (for a constant $^{39}\text{Ar}/^{40}\text{Ar}^*$ in the rock)

(GCA; Yokochi, Sturchio & Purtschert, 2012)

$^{39}\text{Ar}/^{40}\text{Ar}^*$ Chronometer

Details

- Age

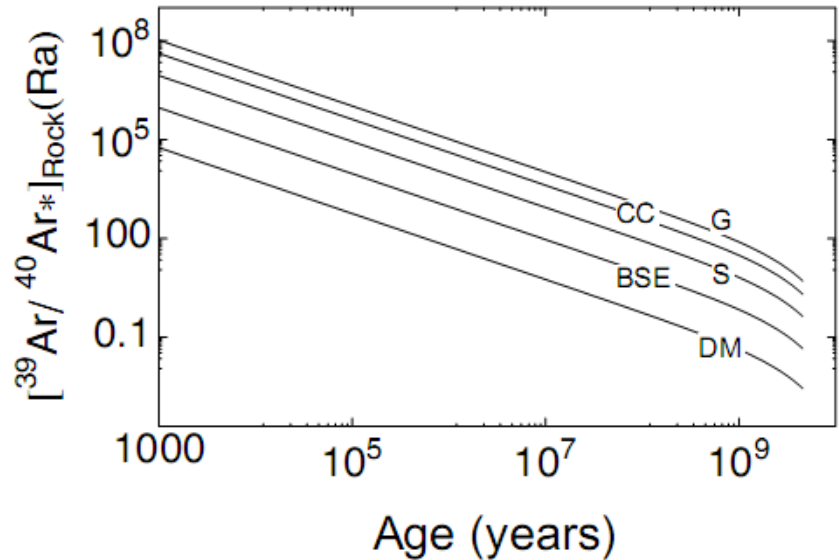
$$t_F \gg \frac{R_C}{R_F} \times \frac{1}{\lambda_{39}}$$

- Time range

- Detectable $^{40}\text{Ar}/^{36}\text{Ar}$ anomaly relative to initial
- Detectable amount of ^{39}Ar
- Up to a few Myr ($\gg 1800$ yrs)

Difficulties+

→ Closed system Estimate



- Deviation associated with gas loss
- Production rate?
- Extent of source rock?

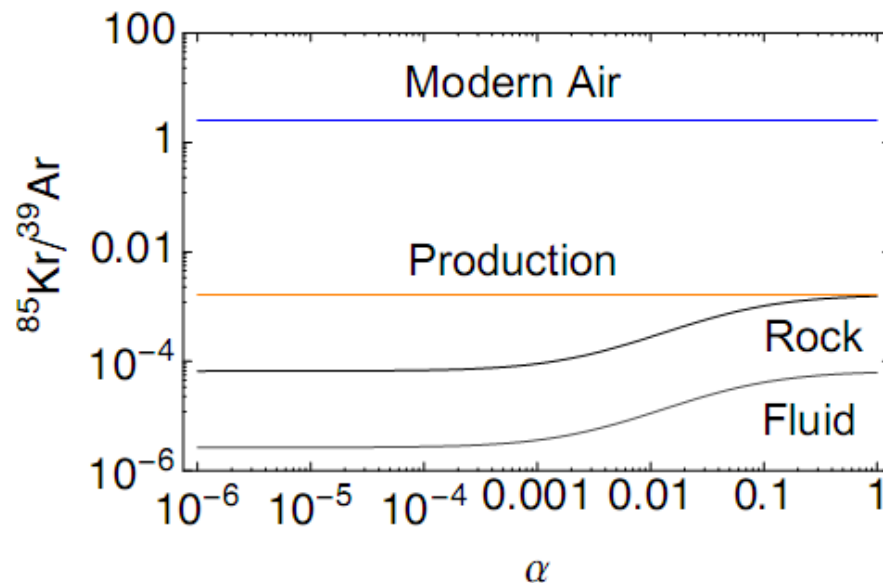
(GCA; Yokochi, Sturchio & Purtschert, 2012)

Krypton-85

Is subsurface ^{85}Kr significant?

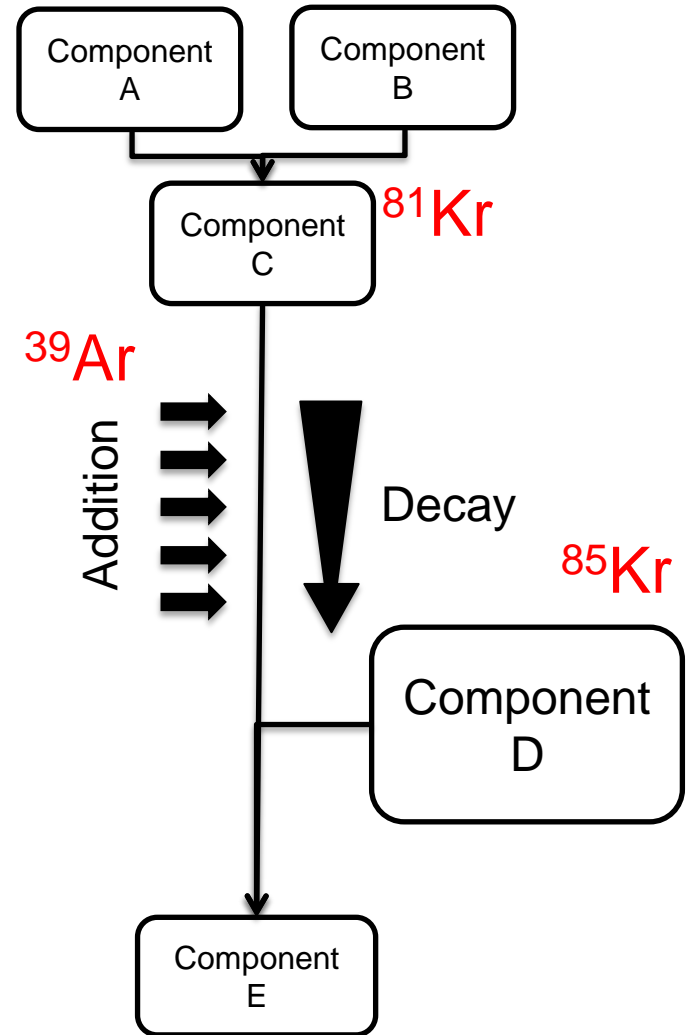
- Production Rates
 - ^{85}Kr : $2.5\text{-}8 \times 10^{-4}$ atoms/g/yr
 - ^{39}Ar : $0.04\text{-}0.9$ atoms/g/yr
- $^{85}\text{Kr}/^{39}\text{Ar}$: $0.001\text{-}0.01$
 - Modern Air \gg Subsurface
 - Surface contamination dominates ^{85}Kr budget
 - Good tracer of shallow contamination

$^{85}\text{Kr}/^{39}\text{Ar}$ ratios



Summary 1: Model Study

- Subsurface-produced ^{39}Ar can serve as chronometer when combined with $^{40}\text{Ar}^*$.
- ^{85}Kr and $^{85}\text{Kr}/^{39}\text{Ar}$ ratio are ideal tracers of modern atmospheric contamination.
- There is no subsurface production of ^{81}Kr .

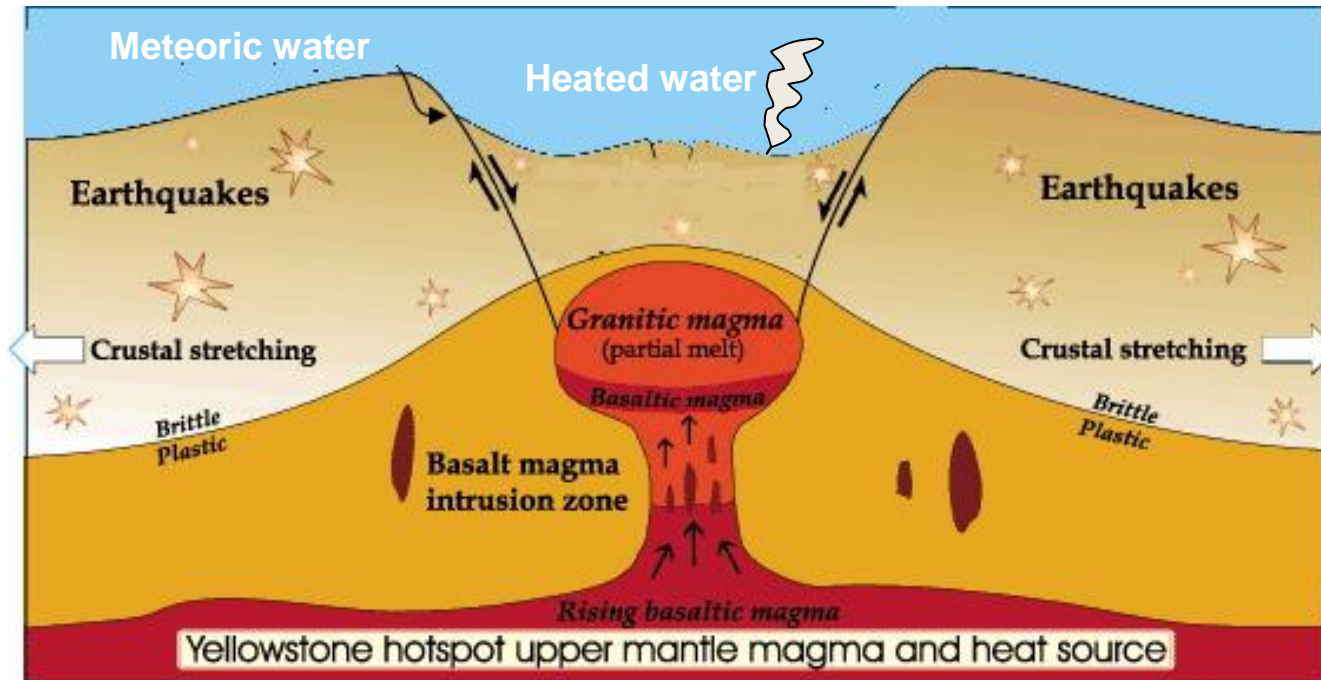




Case Study – Yellowstone National Park



Motivation



- How much time does it take for the meteoric water to go down and come back up?



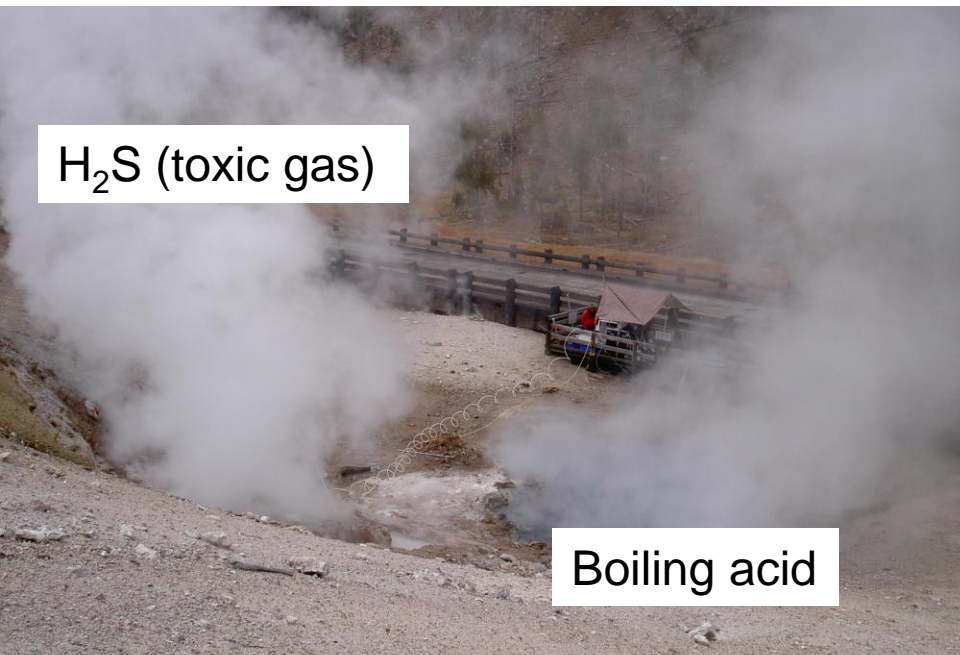
Participants





Frozen

Fieldwork



H_2S (toxic gas)

Boiling acid



Freezing



Beryl Spring



Frying Pan

Sampling Sites



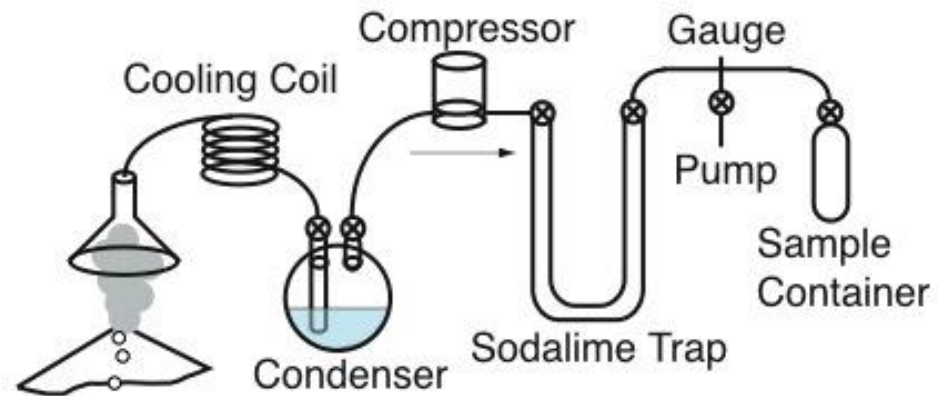
Mud Volcano



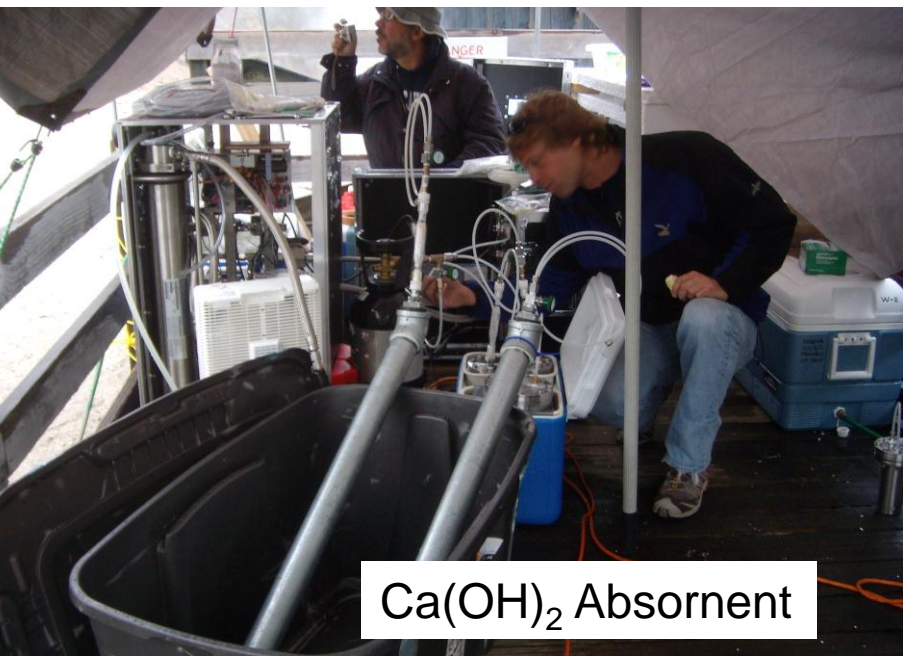
Ojo Caliente



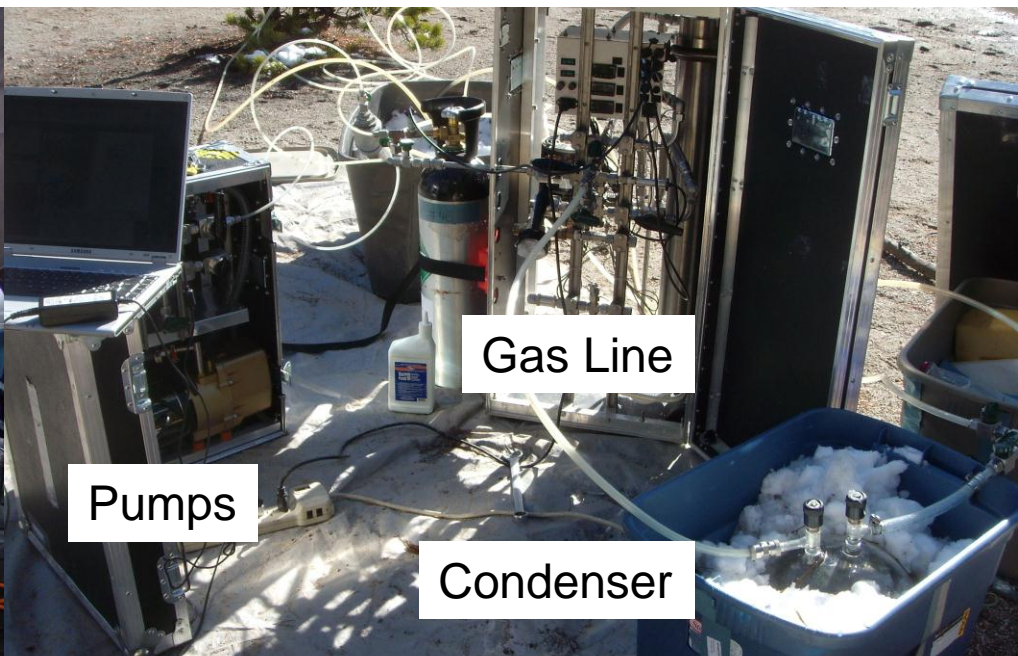
Funnel



Sampling Method



Ca(OH)_2 Absorbent



Pumps

Gas Line

Condenser

Data



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Atom Trap Trace Analysis (ATTA) Report

Report No. 002 Report Date 08 Dec 2011 Project Name Yellowstone

Samples supplied by: Neil Sturchio (UIC), Reika Yokochi (Chicago), Roland Purtschert (Bern)

Samples analyzed by: Wei Jiang, Guo-Min Yang, Peter Mueller, Zheng-Tian Lu (ANL)

Wei Jiang
12-08-2011

ATTA trace No.	Sample No.	Sampling comments	Sampling Date	Size (micro-L)	ATTA Date	⁸⁵ Kr (dpm/cc)	⁸¹ Kr sample / air	ATTA Lab comments
10014	S12	Frying Pan Air, #41	20 Feb 2008	~10	01 Nov 2011	56.2 ± 2.6	1.03 ± 0.057	Local Air
10015	S15	Frying Pan	05 Mar 2008	~10	03 Nov 2011	43.6 ± 2.2	1.03 ± 0.065	
10016	S18	Ojo Caliente	18 Mar 2008	~10	06 Nov 2011	36.9 ± 1.9	1.06 ± 0.063	
10017	S14	Beryl	07 Mar 2008	~3.4	08 Nov 2011	51.1 ± 3.1	0.971 ± 0.090	(air contamination, %99.7 air, P_N2/P_Kr=250 on RGA)
10020	S17	MV1, #42	22 Feb 2008	~1.4	15 Nov 2011	33.4 ± 6.3	1.66 ± 0.33	← Interesting!
10024	S4	MV2, #43	18 Feb 2008	~7	06 Dec 2011	58.9 ± 3.0	0.92 ± 0.061	(Sample cylinder valve leak)

Notes

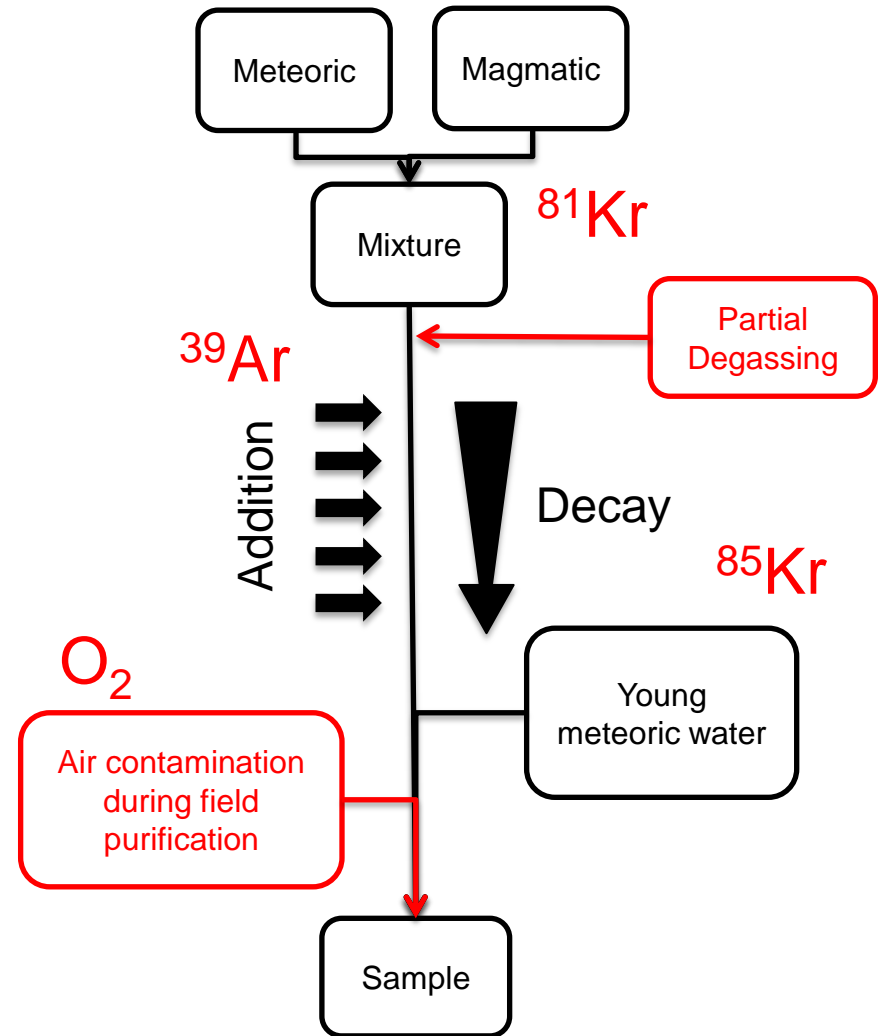
- ⁸⁵Kr ($t_{1/2} = 10.76 \pm 0.02$ yr) abundance is reported in the traditional unit of dpm/cc (decays per minute per cc STP of krypton).
 - Conversion: 100 dpm/cc corresponds to the isotopic abundance of 1.03 ± 0.057
 - The reported ⁸⁵Kr value is as measured on the ATTA apparatus
- ⁸¹Kr ($t_{1/2} = 229 \pm 11$ kyr) abundance is reported as the sample-to-air ratio: $(\text{dpm}/\text{m}^3)_{\text{sample}} / (\text{dpm}/\text{m}^3)_{\text{air}}$.

Need to be evaluated in future work

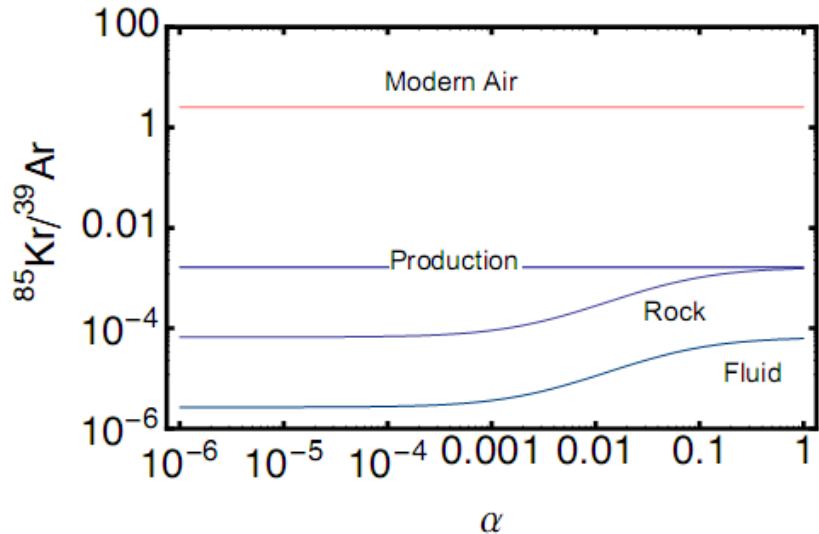
Data Analysis

Other Data

- Chemical composition (UIC)
 - $O_2 > \%$
 - Correction for air contamination during sampling
- Noble gas isotopes (LBNL)
 - $^3\text{He}/^4\text{He}$ and $^{40}\text{Ar}/^{36}\text{Ar}$ variation
 - Partial degassing (Kennedy et al., 1985)
- ^{39}Ar (Purtschert et al., 2009)
 - $^{39}\text{Ar} \gg$ Atmospheric (up to $\times 6$)

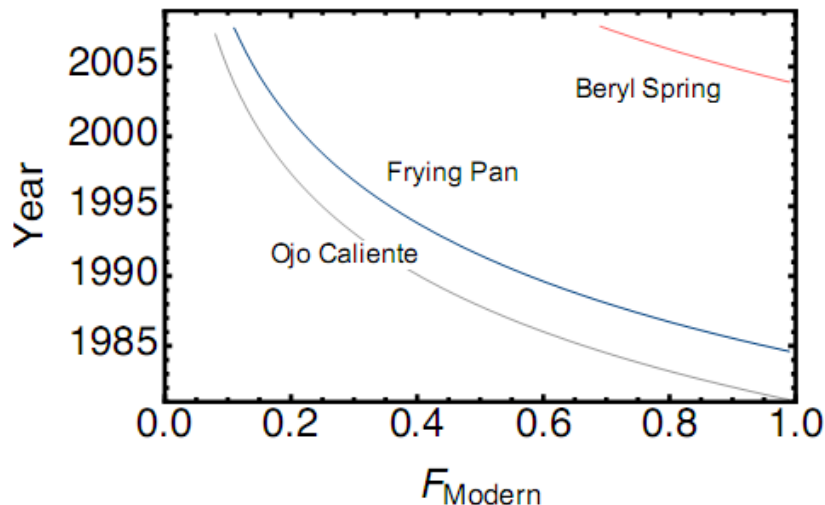


Discussions 1: ^{85}Kr



Fraction of young air In the samples

- Clear contribution of modern ^{85}Kr
- Total: 36-74% Kr contamination during sampling = corrected
- Fraction of young air contribution depends on the age of the air component.
- Lower limit on the fraction of modern ^{81}Kr = **Upper age limit goes up** (Although measured values were atmospheric)



Discussions 2: ^{39}Ar

Reservoir Rocks and Production Rates

- Lava Creek Tuff
Age: 640 kyr
 $^{39}\text{Ar}_{\text{pro}}$: 0.79 atoms/g/yr
 $^{39}\text{Ar}/^{40}\text{Ar}^* = 1.3 \times 10^5 \text{ Ra}$
- Central Plateau Member
Age: 160 kyr
 $^{39}\text{Ar}_{\text{pro}}$: 0.57 atoms/g/yr
 $^{39}\text{Ar}/^{40}\text{Ar}^* = 9.2 \times 10^4 \text{ Ra}$

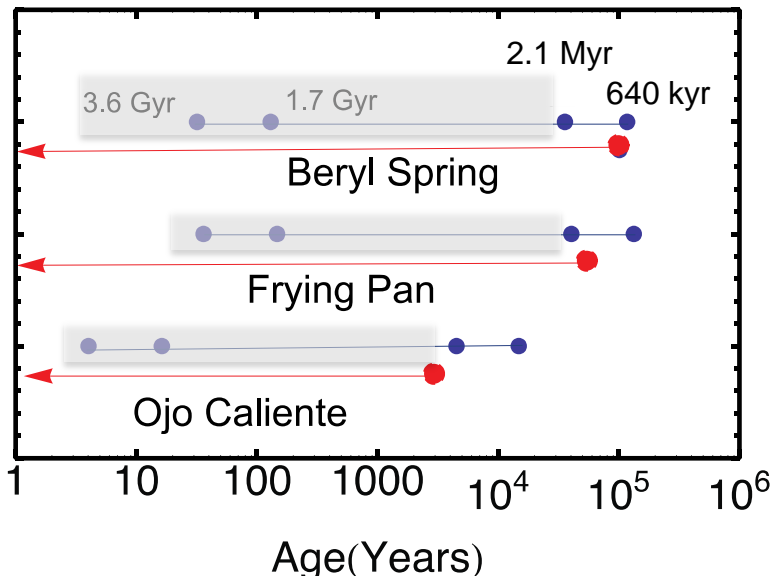
Fluid Data Analysis

$$t_F \gg \frac{R_C}{R_F} \times \frac{1}{\lambda_{39}}$$

- Ages (for closed system)
 - Frying Pan: 130 kyr
 - Beryl Spring: 115 kyr
 - Ojo Caliente: 16 kyr
- External flux?
 - Ar from older reservoir
=>younger age

Summary 2: YNP

Chronological Constraints



- Geothermal activity over >100 kyr (³⁹Ar age) if noble gas source is the reservoir rock
- ⁸¹Kr age: Consistent with Ar age

Discussions

- The chronometer will improve if another parameter determines
 - The fraction (or age) of young meteoric component (⁸¹Kr)
 - The source of radiogenic/nucleogenic isotopes (³⁹Ar) (reservoir rock vs. Pre-Cambrian basement)
- Systematic, multi-component study is essential for interpreting the complex system

Conclusions

- Nucleogenic ^{39}Ar , an obstacle of ventilation age dating, can serve as a chronometer when combined with $^{40}\text{Ar}^*$.
- Isotopic abundances of noble gas radionuclides in geothermal gases from Yellowstone National Park were analyzed for the first time, suggesting long water residence time in the crust
- Systematic, multi-component study is essential for interpreting the complex system
- $^{39}\text{Ar}/^{40}\text{Ar}^*$ chronometer is applicable for old groundwater dating as well as tracing the time scale of fluid migration in a variety of scenarios including analogue and pilot studies of CO_2 sequestration and the formation of petroleum or natural gas reservoirs.
- Looking forward to new data!